

# Finite Elements — Motivation Lecture

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October 18, 2021

Organizational Stuff

Partial Differential Equations are Ubiquitous

Preview: The Finite Element Method

DUNE Software

## Organizational Staff

- ▶ Lecturer: Peter Bastian  
Office: INF 205, room 1/401  
email: `peter.bastian@iwr.uni-heidelberg.de`
- ▶ Lecture: Mi + Fr 9-11 SR C  
Exercise: Tue 14-16 SR 11
- ▶ Lecture homepage:  
`https://conan.iwr.uni-heidelberg.de/teaching/finiteelements\_ws2021/`
- ▶ Lecture notes available from the homepage
- ▶ Moodle page `https://moodle.uni-heidelberg.de/course/view.php?id=10404`  
inscription should be possible without key
- ▶ Lecture will be recorded (writing and audio) and put on moodle

- ▶ Excercises Tue 14-16 in SR 11 organized by Michal Tóth  
Office: INF 205, room 01.224  
email: `michal.toth@iwr.uni-heidelberg.de`
- ▶ Registration to excercises via MÜSLI system:  
`https://muesli.mathi.uni-heidelberg.de/lecture/view/1420`
- ▶ There will be theoretical and practical excercises
- ▶ Practical excercises are important! They will be based on the software DUNE: `www.dune-project.org`  
*You will need some UNIX environment (Linux/MacOS)*
- ▶ Scheme:
  - ▶ Exercises can be done in groups of 2...3
  - ▶ Exercise given out / handed in Monday evening
  - ▶ Discussion of submitted exercises on Tuesday
  - ▶ You self-grade your exercise: how many points is it worth?
  - ▶ Random selection of presenting groups (a group may be selected even if no member is present!)
- ▶ Exercises start Tue, October 26

- ▶ Written exam (Klausur) at the end of the semester
- ▶ Date: February 18, 2022
- ▶ Requirements:

- ▶ Be interactive! Ask questions!
- ▶ There are no dumb questions!
- ▶ Do not miss the practical exercises. Polish your C++ knowledge!

## Partial Differential Equations are Ubiquitous

- ▶ Calculus was invented for (partial) differential equations!
- ▶ E.g. to express conservation of mass, momentum and energy in quantitative form
- ▶ Famous examples are:
  - ▶ Poisson (electrostatics, gravity) 1800
  - ▶ Euler (inviscid flow) 1757
  - ▶ Navier-Stokes (viscous flow) 1822/1845
  - ▶ Maxwell (electrodynamics) 1864
  - ▶ Einstein (general relativity) 1915
- ▶ Solutions in practical situation only with modern (super) computers!



Poisson



Euler



Navier



Stokes



Maxwell



Einstein

- ▶ This is a numerics class, but . . .
- ▶ In order to judge whether a numerically computed solution is reasonable, it is good to have an understanding of the underlying application problem
- ▶ In the first part of the lecture we will look at deriving some models
- ▶ Let us look at some examples for motivation!

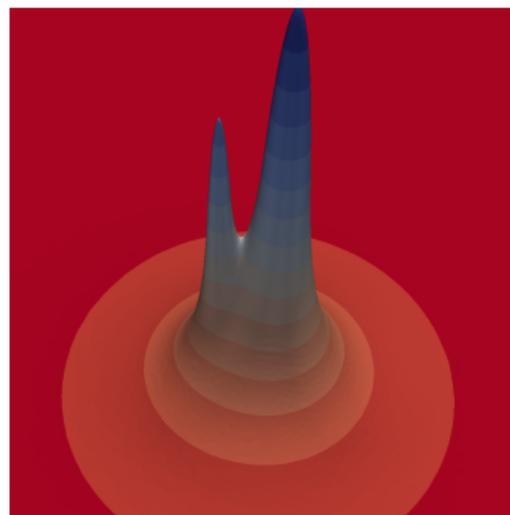
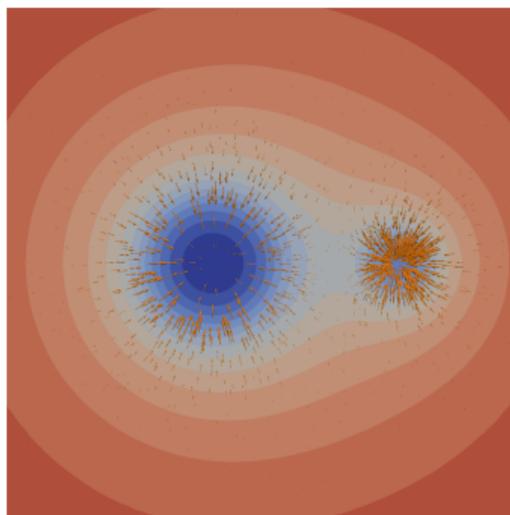
## Gravitational Potential (Poisson Equation)

Find function  $\Psi(x) : \Omega \rightarrow \mathbb{R}$ ,  $\Omega = \mathbb{R}^3$  such that:

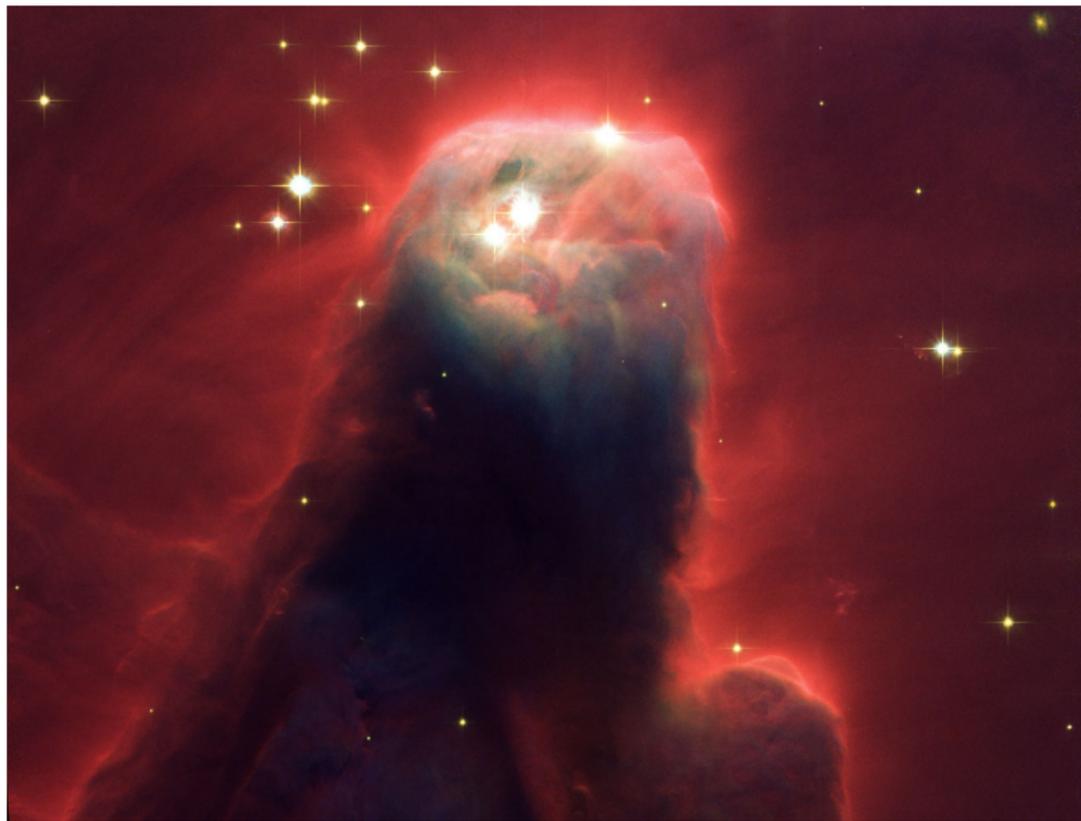
$$\partial_{x_1 x_1} \Psi(x) + \partial_{x_2 x_2} \Psi(x) + \partial_{x_3 x_3} \Psi(x) = \nabla \cdot \nabla \Psi(x) = \Delta \Psi(x) = 4\pi G \rho(x)$$

$G$ : gravitational constant,  $\rho$ : mass density in  $\text{kg}/\text{m}^3$

Force acting on small point mass  $m$  at point  $x$ :  $F(x) = -m \nabla \Psi(x)$



# Star Formation



Cone nebula from <http://www.spacetelescope.org/images/heic0206c/>

Euler equations of gas dynamics:

$$\partial_t \rho + \nabla \cdot (\rho \mathbf{v}) = 0 \quad (\text{mass conservation})$$

$$\partial_t (\rho \mathbf{v}) + \nabla \cdot (\rho \mathbf{v} \mathbf{v}^T + p \mathbf{I}) = -\rho \nabla \Psi \quad (\text{momentum conservation})$$

$$\partial_t e + \nabla \cdot ((e + p) \mathbf{v}) = -\rho \nabla \Psi \cdot \mathbf{v} \quad (\text{energy conservation})$$

$$\Delta \Psi = 4\pi G \rho \quad (\text{gravitational potential})$$

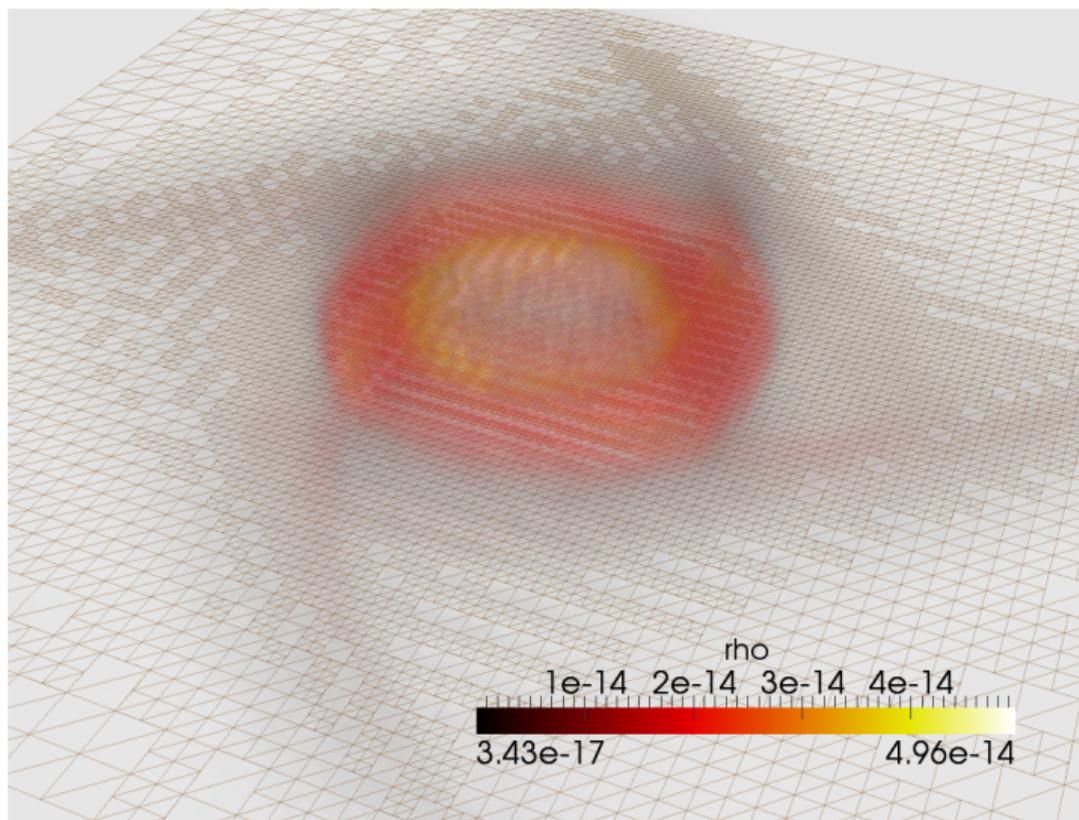
Constitutive relation:  $p = (\gamma - 1)(e - \rho \|\mathbf{v}\|^2/2)$

Plus the Poisson equation ...

More elaborate model requires radiation transfer, better constitutive relations, friction, ...

Nonlinear system of partial differential equations

# Star Formation: Numerical Simulation



(Diploma thesis of Marvin Tegeler, 2011)

(Incompressible) Navier-Stokes Equations:

$$\nabla \cdot \mathbf{v} = 0 \quad (\text{mass conservation})$$

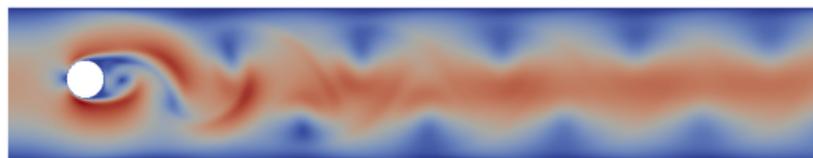
$$\partial_t \mathbf{v} + \nabla \cdot (\mathbf{v}\mathbf{v}^T) - \nu \Delta \mathbf{v} + \nabla p = \mathbf{f} \quad (\text{momentum conservation})$$

- ▶  $\rho$  is independent of pressure
- ▶ No compression work, isothermal situation
- ▶ Pressure is independent variable
- ▶ Existence of solutions is Millenium Prize Problem (in 3d for general data)

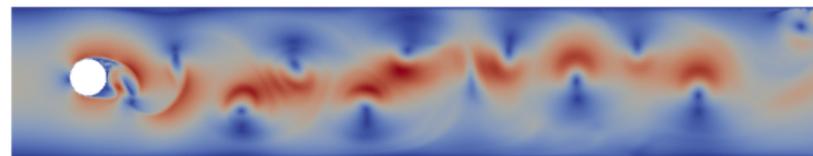
# Von Karman Vortex Street



Re 20 (laminar)

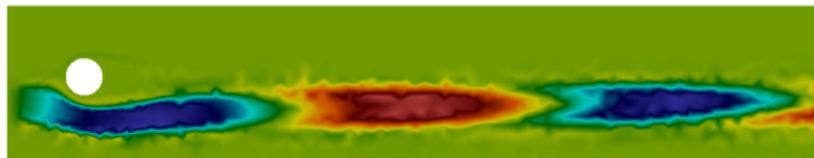


Re 200 (periodic)

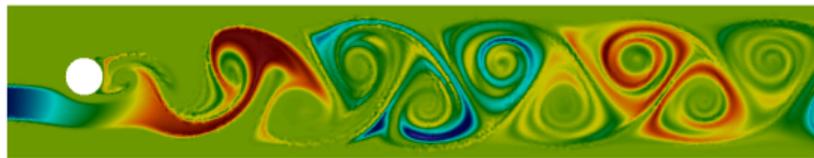


Re 1500 (turbulent)

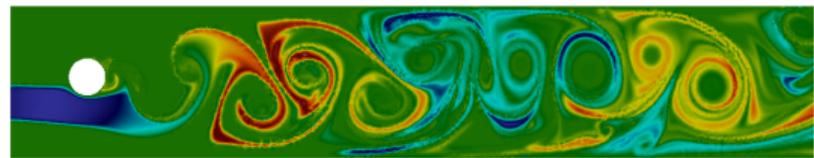
# Von Karman Vortex Street



Re 20



Re 200



Re 1500

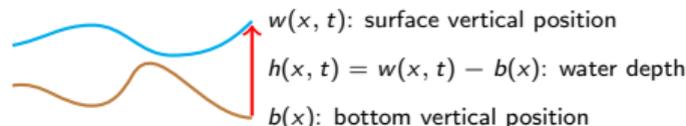
- ▶ Starting point: the incompressible Navier-Stokes equations

$$\begin{aligned}\partial_t(\rho v) + \nabla \cdot (\rho v v^T) - \nabla \cdot \sigma(v, p) &= \rho f \\ \nabla \cdot (\rho v) &= 0\end{aligned}$$

with  $v$  velocity,  $p$  pressure,  $\rho$  (const.) density,  $\sigma$  given by

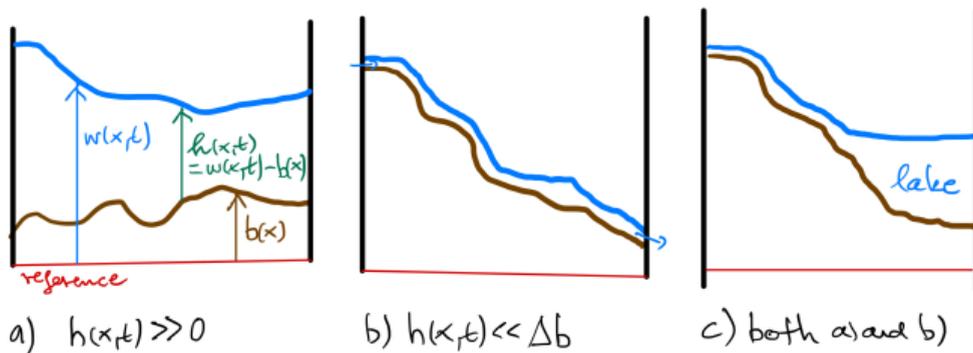
$$\sigma(v, p) = 2\mu\epsilon(v) - pl \quad (\text{stress}), \quad \epsilon(v) = \frac{1}{2}(\nabla v + (\nabla v)^T) \quad (\text{strain rate})$$

- ▶ Free surface, no breaking waves  $\rightsquigarrow$  time-dependent domain



- ▶ Boundary conditions:

- ▶ No-slip boundary condition:  $v(x, t) = 0$  (bottom, lateral sides)
- ▶ Navier slip condition  $v \cdot n_b = 0$ ,  $(n_b \times (\sigma(v, p)n_b)) \times n_b + \frac{\beta}{\|n_b\|^3} v = 0$
- ▶ No flow at free surface:  $v(x, t) \cdot n_w(x, t) = \frac{\partial w}{\partial t}(x, t)$ ,  $w(x, t)$  surface pos.



- ▶ Extend of fluid domain much larger in horizontal than in vertical direction
- ▶ a) ocean flow, b) and c) land surface flow
- ▶ Dimension reduction reduces computational work

$$\partial_t(\rho h(\hat{x}, t)) + \sum_{j=1}^2 \partial_j(\rho h(\hat{x}, t) \bar{v}_j(\hat{x}, t)) = 0,$$

$$\begin{aligned} \partial_t(\rho h \bar{v}_i) + \sum_{j=1}^2 \partial_j(\rho h \bar{v}_i \bar{v}_j) + \rho g \frac{\partial w}{\partial x_i}(\hat{x}, t) h(\hat{x}, t) \\ - 2\mu [(\epsilon(\hat{x}, w(\hat{x}, t), t) n_w(\hat{x}, t))_i + (\epsilon(\hat{x}, b(\hat{x}, t), t) n_b(\hat{x}))_i] = 0. \quad (i = 1, 2) \end{aligned}$$

- ▶ First-order hyperbolic system for water height  $h$  and vertically averaged horizontal velocity  $\bar{v}$

$$\bar{v}_i(x_1, x_2, t) = \frac{1}{h(x_1, x_2, t)} \int_{b(x_1, x_2)}^{w(x_1, x_2, t)} v_i(x_1, x_2, x_3, t) dx_3 \quad i = 1, 2.$$

- ▶ Derived rigorously from Navier-Stokes under very few assumptions:
  - ▶  $v_3$  is very small  $\rightsquigarrow$  hydrostatic pressure assumption
  - ▶ Velocities do not deviate much from their average  $\rightsquigarrow \overline{v_i v_j} \approx \bar{v}_i \bar{v}_j$
  - ▶ No internal friction (but surface and bottom friction)

- ▶ Model friction, e.g.  $(\epsilon(\hat{x}, b(\hat{x}), t)n_b(\hat{x}))_i = \alpha \bar{v}_i$  (Navier slip)
- ▶ One-dimensional shallow water equations: St. Venant equations
- ▶ Diffusive wave approximation

$$\frac{\partial(\rho h)}{\partial t} - \nabla \cdot (c_{dw}\rho h^2(\hat{x}, t)\nabla(h + b)) = 0.$$

- ▶ Momentum equation: keep only gravity and bottom friction
  - ▶ Employ Navier slip condition
  - ▶ Insert into mass conservation
  - ▶ Nonlinear diffusion equation, other nonlinearities are used below
- ▶ Kinematic wave approximation

$$\frac{\partial(\rho h)}{\partial t} - \nabla \cdot (c_{kw}\rho h^2(\hat{x}, t)\nabla b(\hat{x})) = 0.$$

- ▶ In addition assume  $\partial_i w \approx \partial_i b$  (no lake at rest ...)
- ▶ Nonlinear first-order hyperbolic, similar to Burger's equation

# (Shallow) Subsurface Flow

- ▶ Flow in fully saturated porous medium (groundwater flow equation):

$$\nabla \cdot v(x, t) = f \quad (\text{mass cons.}), \quad v(x, t) = -\frac{K(x)}{\mu} (\nabla p(x, t) - \rho g) \quad (\text{Darcy})$$

$K$  permeability,  $g$  gravity vector

- ▶ Can be derived from Stokes equations
- ▶ Confined aquifer: elliptic PDE
- ▶ Unconfined aquifer
  - ▶ With capillary effects: two-phase flow, Richards equation
  - ▶ Without capillary effects: Groundwater flow with free surface
- ▶ Groundwater flow with free surface and shallow water assumption:

$$\partial_t(\phi u) - \nabla \cdot \left( K \frac{\rho g}{\mu} h \nabla (h + b) \right) = f$$

$\phi$  porosity,  $b$  bathymetry

$$\partial_t w_s - \nabla \cdot \left( \frac{1}{n(x)} \frac{(w_s - b_s(x))^\alpha}{\|\nabla w_s\|^{1-\gamma}} \nabla w_s \right) = f_s(x, t) - q(x, w_s, w_a)$$

$$\phi \partial_t (\min(w_a, b_s) - b_a) - \nabla \cdot (k(x)(w_a - b_a) \nabla w_a) = f_a(x, t) + q(x, w_s, w_a)$$

- ▶ Unknown functions:  $w_s$  surface water level,  $w_a$  groundwater level
- ▶  $n$  Manning's number,  $0 < \alpha \leq 2$ ,  $0 < \gamma \leq 1$
- ▶  $w_a < b_s$ : unconfined aquifer,  $w_a \geq b_s$ : confined aquifer
- ▶ Exchange term (infiltration - exfiltration)

$$q(x, w_s, w_a) = L_i \frac{\max(w_s - b_s(x), 0)}{C + \max(w_s - b_s(x), 0)} \max(w_s - w_a, 0) - L_e \max(w_a - w_s, 0)$$

Instantaneous transfer from surface to groundwater

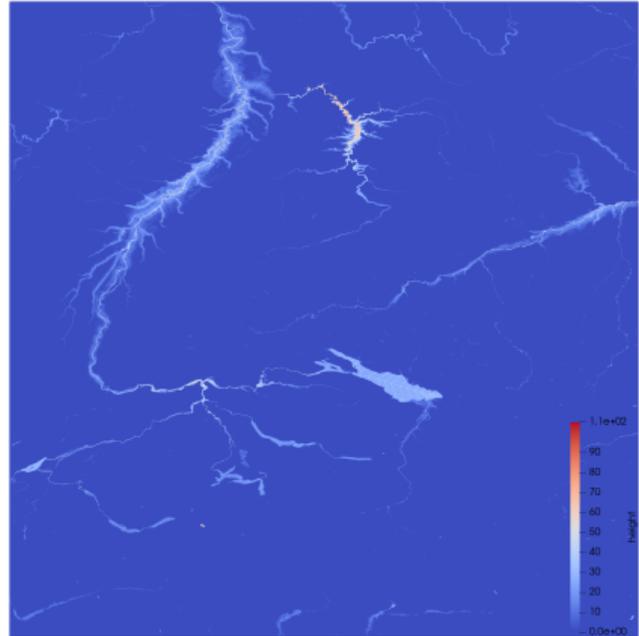
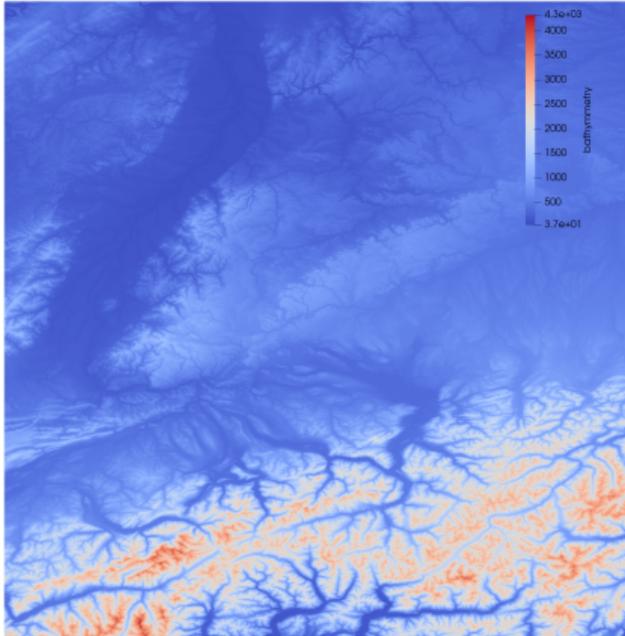
- ▶ Can model rivers, lakes, surface flow, groundwater flow

### Setup:

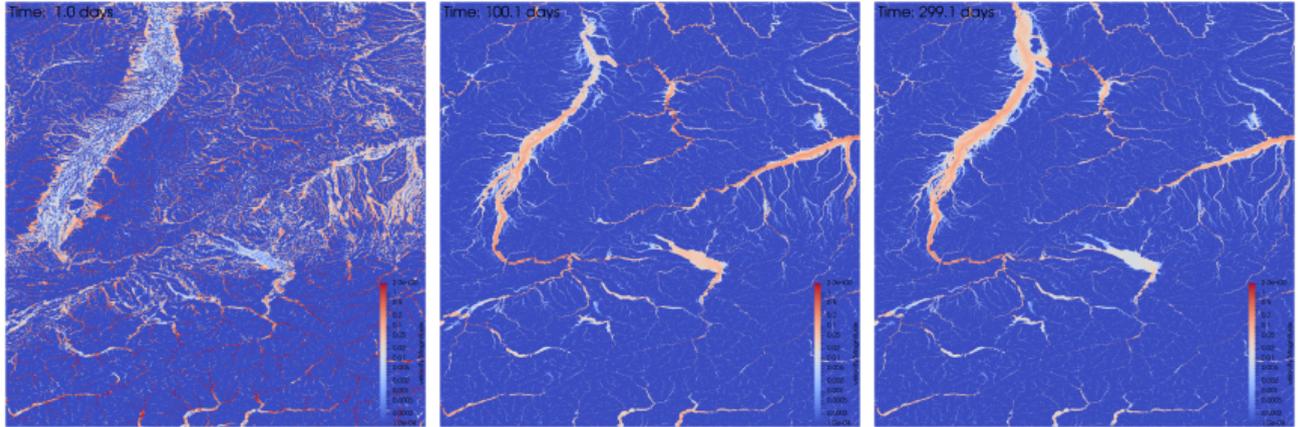
- ▶ Only surface flow model with  $n = 1$ ,  $\alpha = 1$ ,  $\gamma = 1/2$
- ▶ “Hydrologically conditioned” digital elevation model (DEM) from Hydrosheds<sup>1</sup> at 90m resolution
- ▶ 4800<sup>2</sup> cells
- ▶ Constant forcing average rainfall (2 liter per  $m^2$  and day)

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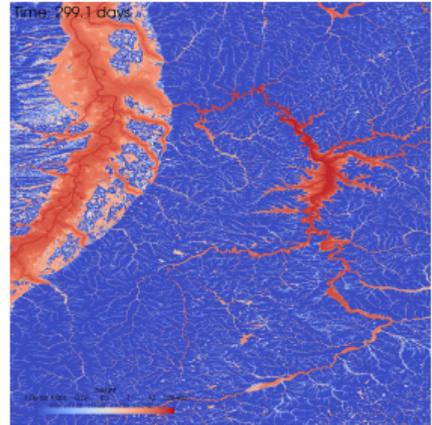
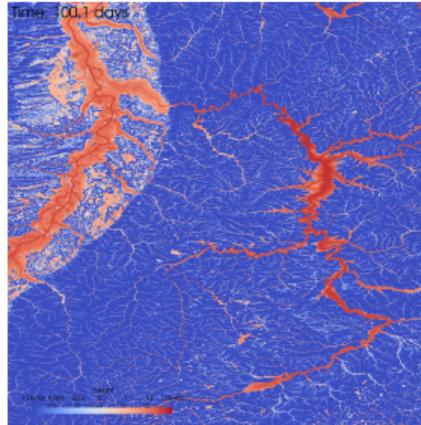
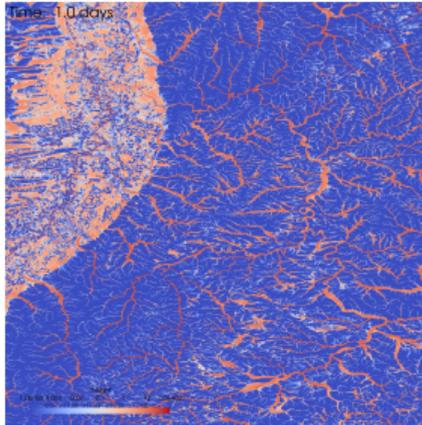
<sup>1</sup><https://www.hydrosheds.org/>



- ▶ Left: Hydrologically conditioned DEM from [www.hydrosheds.org](http://www.hydrosheds.org)
- ▶ Right: Water height after 300 days with initial height 1m



- ▶ Left: flow velocity after 1 day, initial height 1m
- ▶ Middle: flow velocity after 100 days, initial height 1m
- ▶ Right: flow velocity after 300 days, initial height 1m



- ▶ Focus on Neckar river, water height
- ▶ 1m initial height, 1, 100, 300 days

# The Parameter Problem

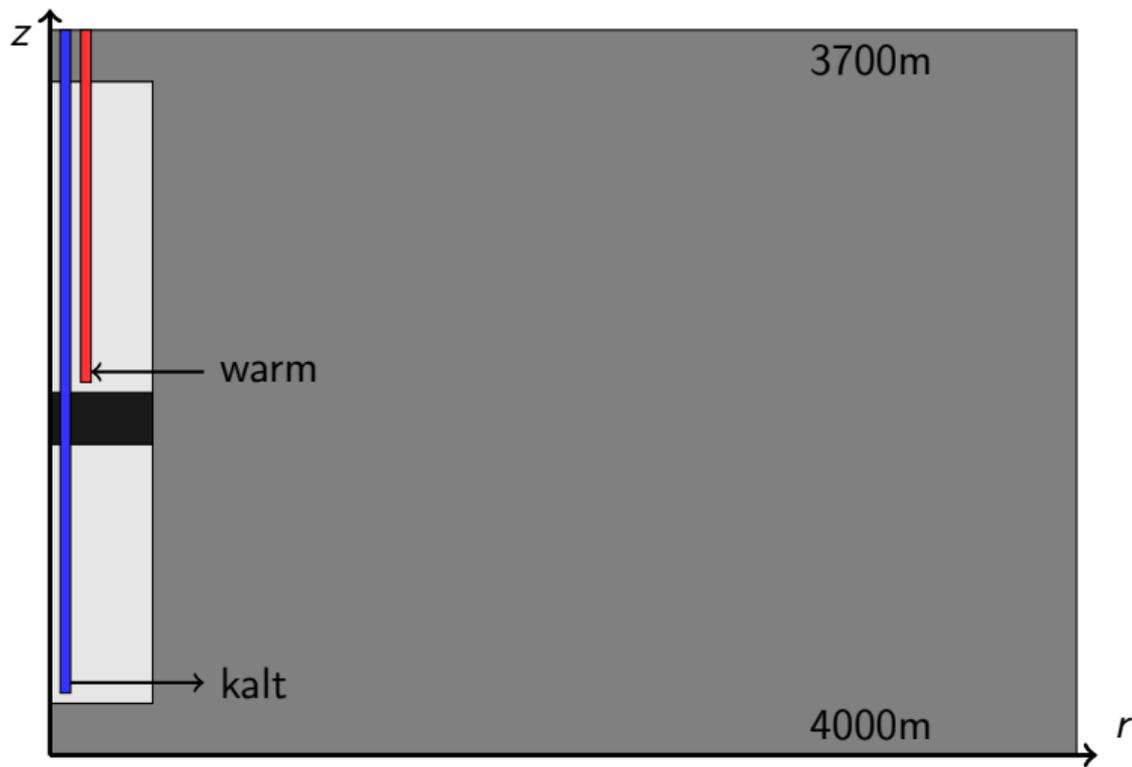
Such high fidelity models need many spatially distributed parameters:

- ▶  $n, \alpha, \gamma$  parameters in surface flow model
  - ▶ depend on land cover
- ▶  $\phi, K$  porosity and permeability in groundwater flow
  - ▶ difficult, permeability varies by orders of magnitude
- ▶  $L_i, L_e$  in exchange terms
  - ▶ difficult, interface free flow / porous medium difficult to model
- ▶  $b_s, b_a$  bathymetries
  - ▶  $b_s$  needs to be “hydrologically conditioned”,  $b_a$  much more uncertain
- ▶  $f_s, f_a$  precipitation, groundwater pumping
  - ▶ precipitation available, but at coarse resolution

Data that can be used for parameter estimation:

- ▶ River levels and discharge
- ▶ Groundwater levels

# Geothermal Power Plant



Coupled system for water flow and heat transport:

$$\partial_t(\phi\rho_w) + \nabla \cdot \{\rho_w u\} = f \quad (\text{mass conservation})$$

$$u = \frac{k}{\mu}(\nabla p - \rho_w g) \quad (\text{Darcy's law})$$

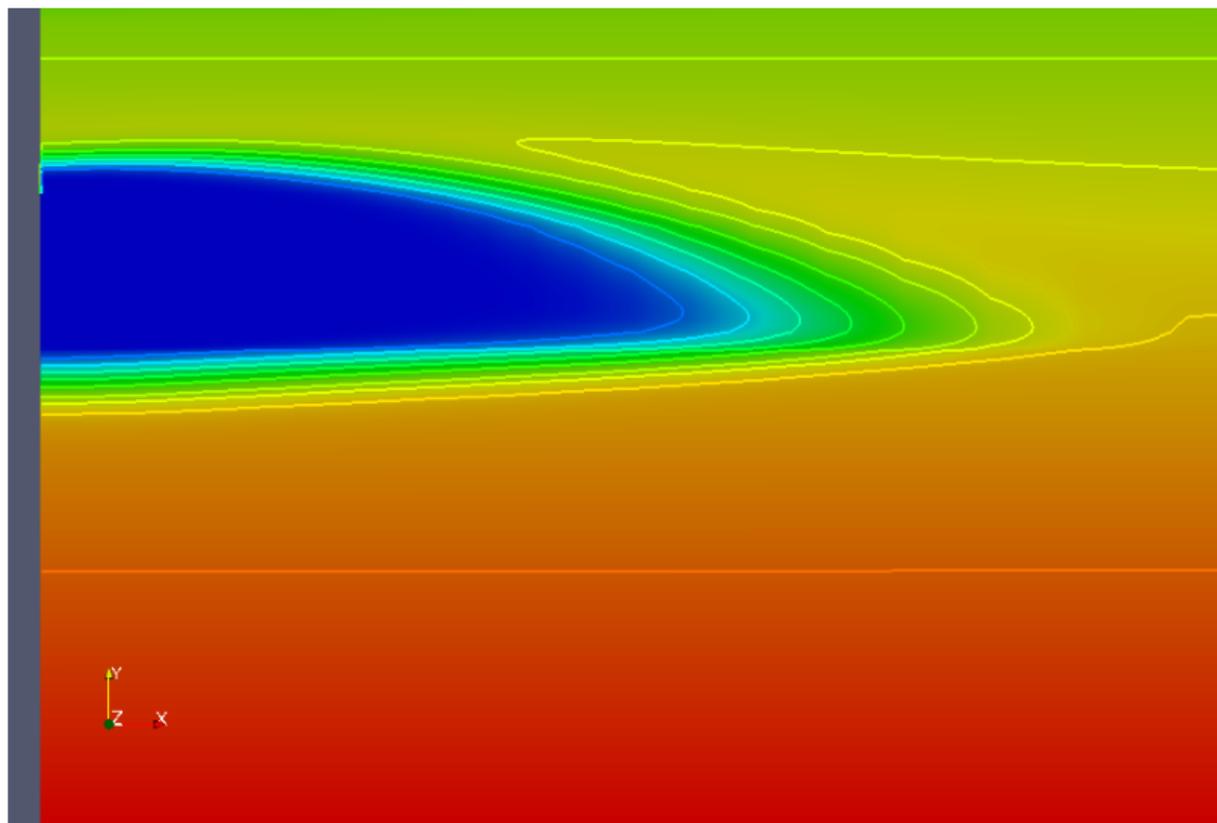
$$\partial_t(c_e \rho_e T) + \nabla \cdot q = g \quad (\text{energy conservation})$$

$$q = c_w \rho_w u T - \lambda \nabla T \quad (\text{heat flux})$$

Nonlinearity:  $\rho_w(T)$ ,  $\rho_e(T)$ ,  $\mu(T)$

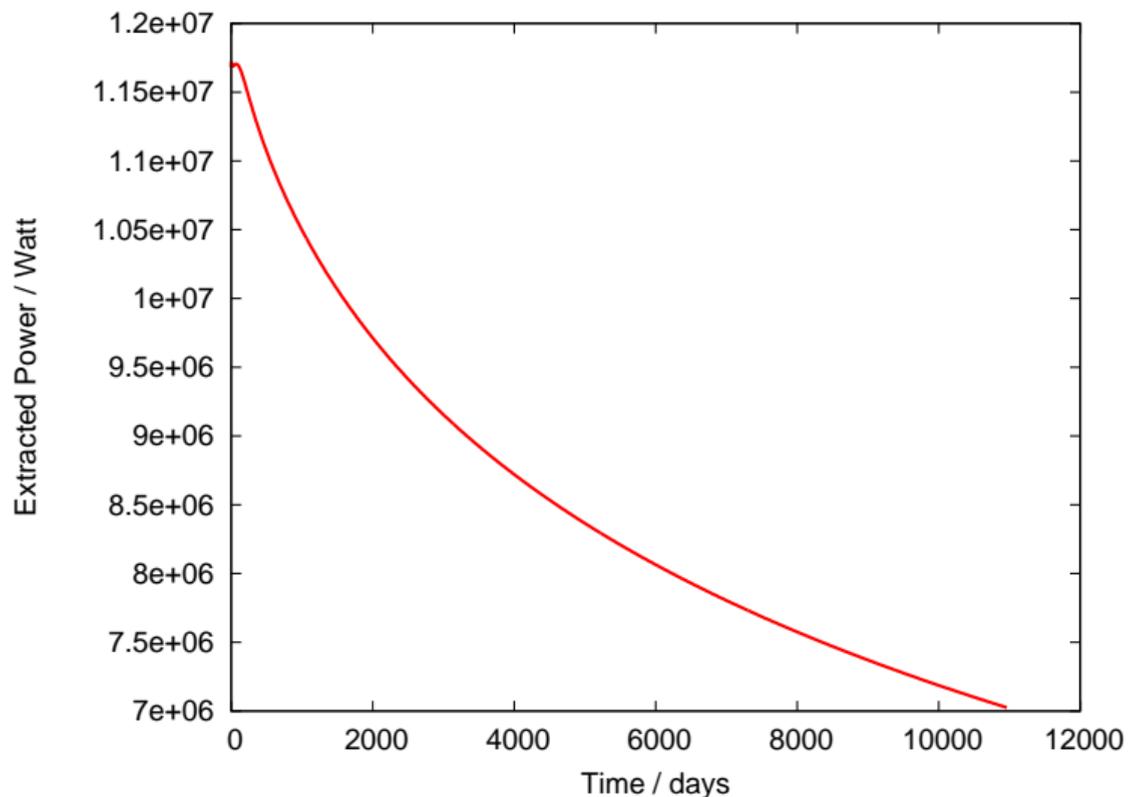
Permeability  $k(x)$  :  $10^{-7}$  in well,  $10^{-16}$  in plug

Space and time scales:  $R=15$  km,  $r_b=14$  cm, flow speed 0.3 m/s in well, power extraction: decades



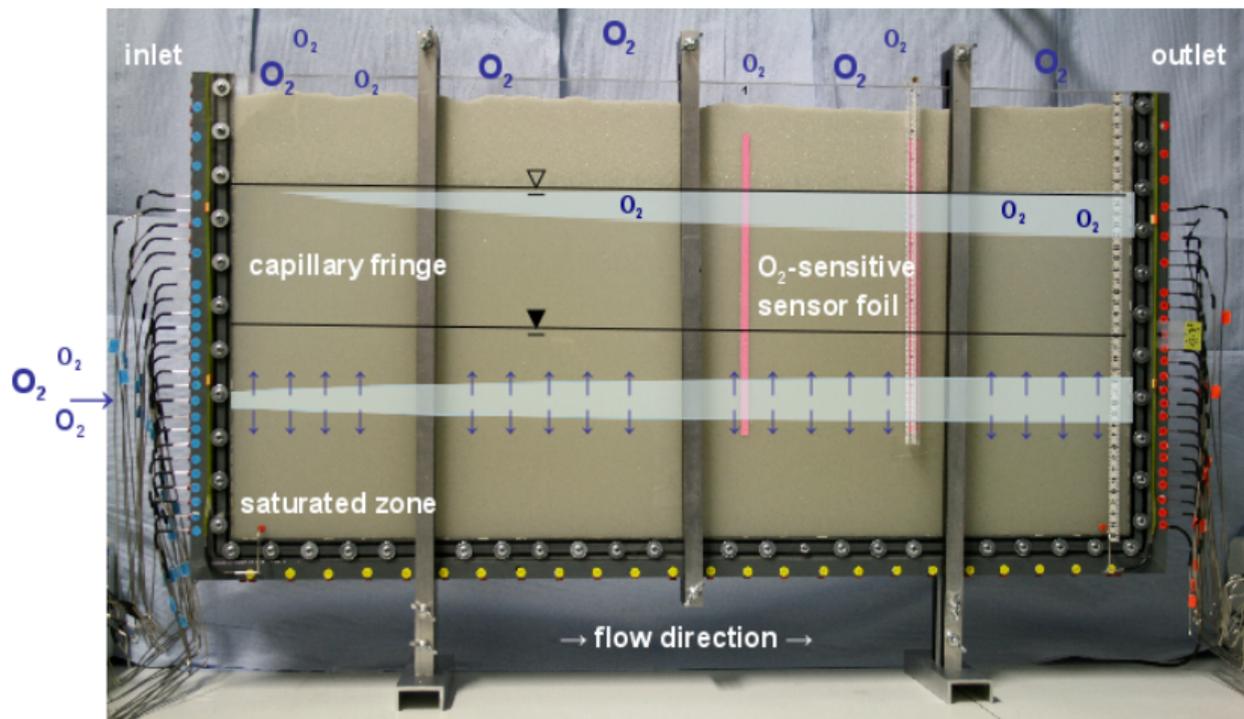
Temperature after 30 years of operation

# Geothermal Power Plant: Results



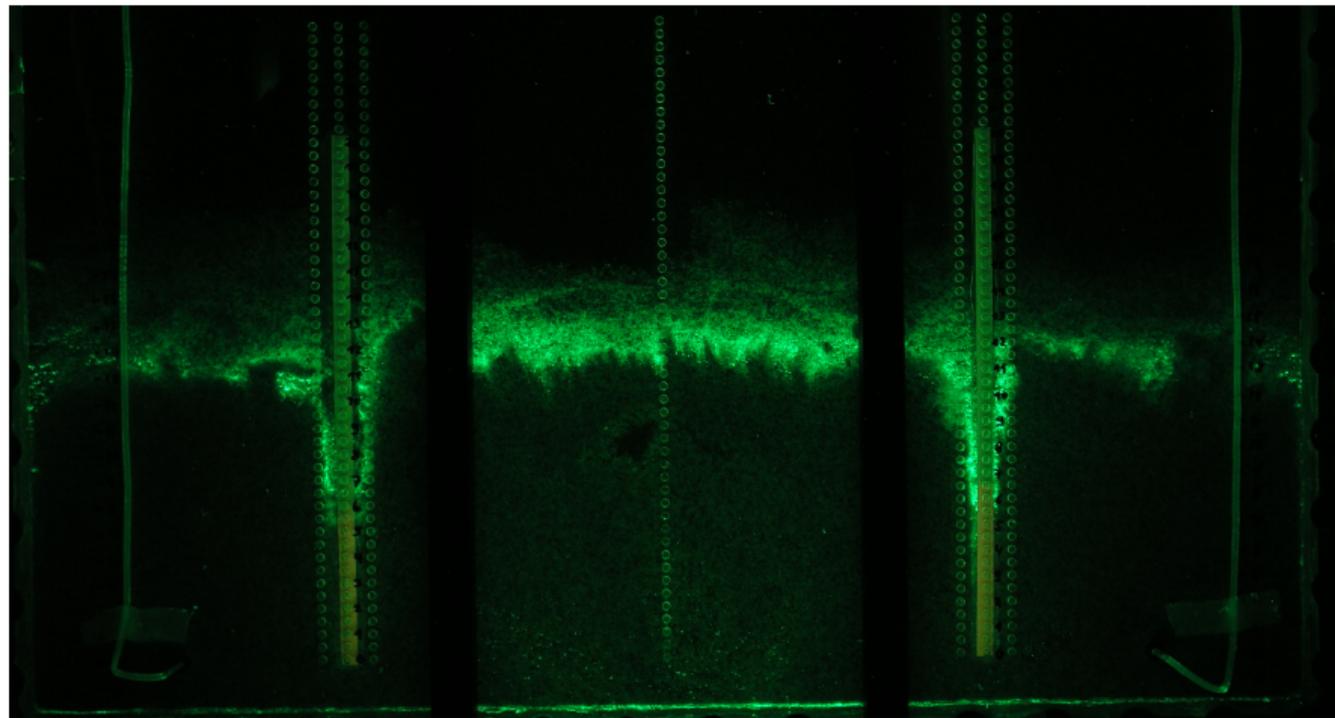
Extracted power over time

# Bacterial Growth and Transport in Capillary Fringe



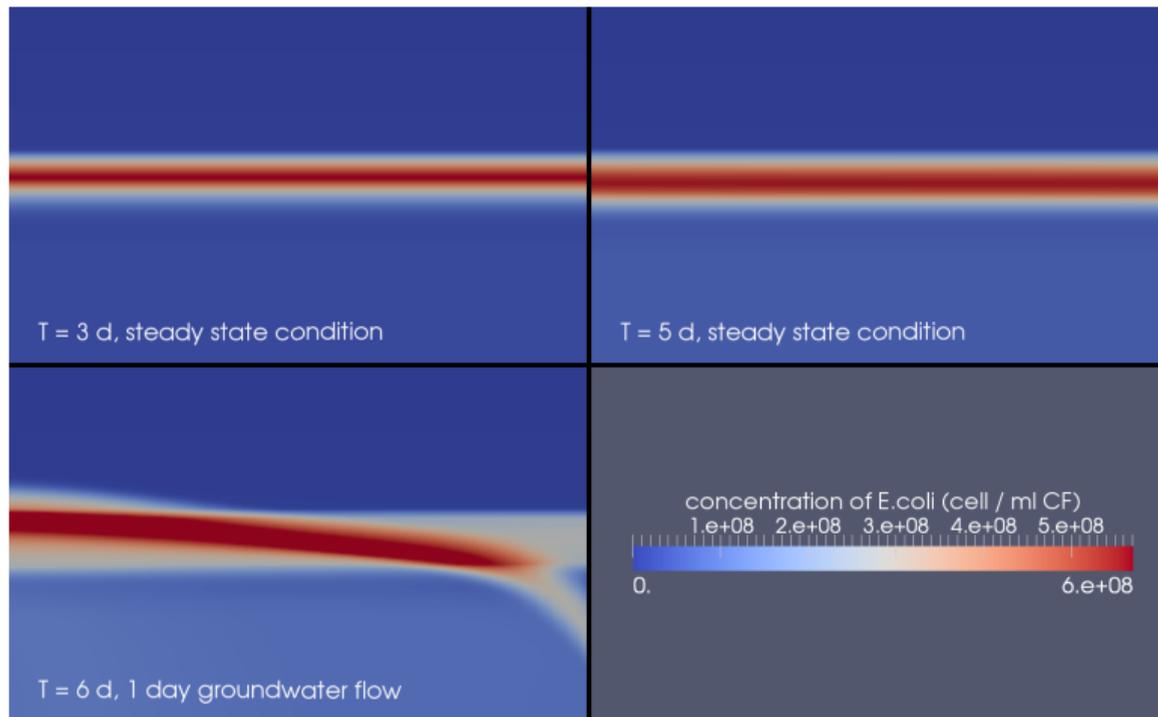
DFG Research Group 831 DyCap, Experiment by C. Haberer, Tübingen

# Bacterial Growth and Transport in Capillary Fringe



Experiment by Daniel Jost, KIT, Karlsruhe

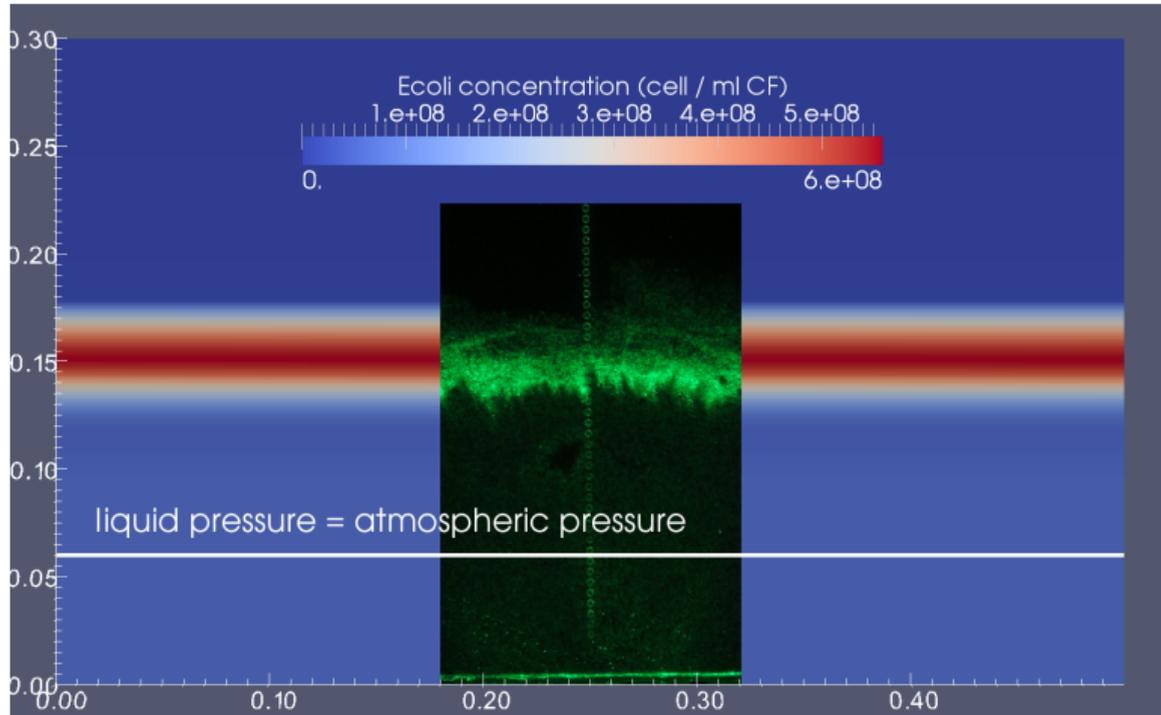
# Reactive Multiphase Simulation



Unknowns: pressure, saturation, bacteria concentration, carbon concentration, oxygen concentration

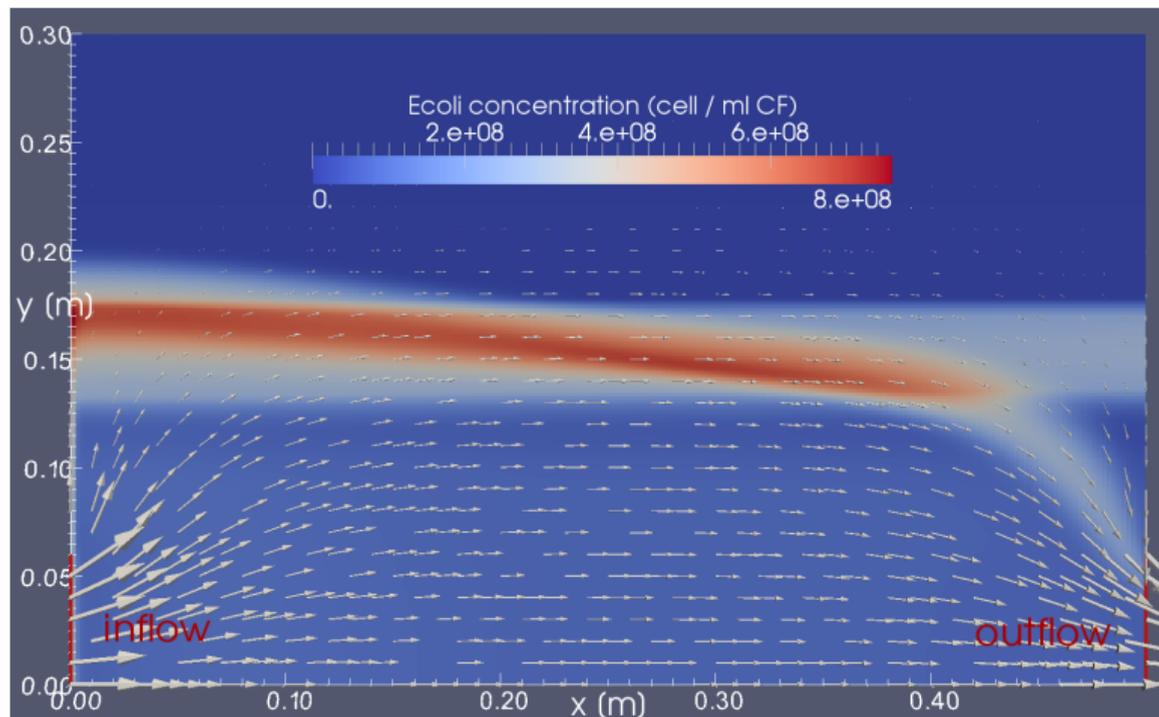
Simulation by Pavel Hron

# Reactive Multiphase Simulation



Simulation by Pavel Hron

# Reactive Multiphase Simulation



Simulation by Pavel Hron

(Macroscopic)Maxwell equations:

$$\nabla \times E = -\partial_t B \quad (\text{Faraday})$$

$$\nabla \times H = j + \partial_t D \quad (\text{Ampère})$$

$$\nabla \cdot D = \rho \quad (\text{Gauß})$$

$$\nabla \cdot B = 0 \quad (\text{Gauß for magnetic field})$$

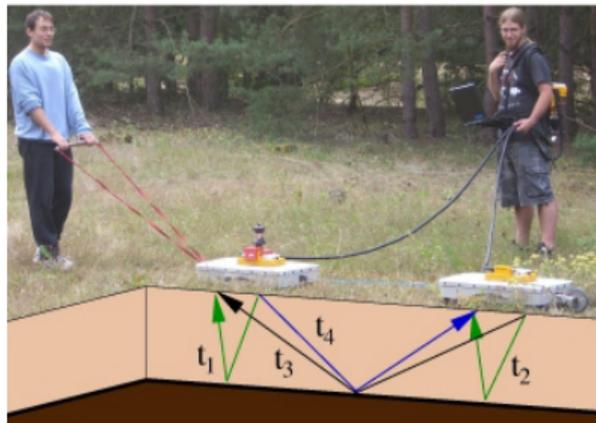
Constitutive relations:

$$D = \epsilon_0 E + P \quad (D: \text{electric displacement field, } P: \text{polarization})$$

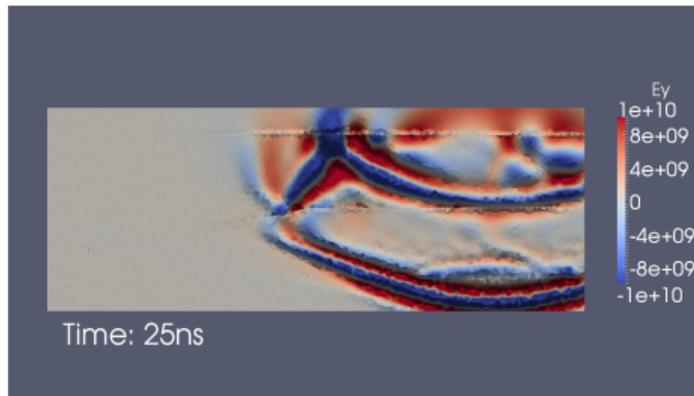
$$B = \mu_0(H + M) \quad (H: \text{magnetizing field, } M: \text{magnetization})$$

Linear, first-order hyperbolic system

# Application: Geo-radar



Soil physics group Heidelberg



Simulation: Jorrit Fahlke

- ▶ Poisson equation: gravity, electrostatics (**elliptic type**)

$$\begin{aligned}\Delta u &= f && \text{in } \Omega \\ u &= g && \text{on } \Gamma_D \subseteq \partial\Omega \\ \nabla u \cdot \nu &= j && \text{on } \Gamma_N = \subseteq \partial\Omega \setminus \Gamma_D\end{aligned}$$

- ▶ Heat equation (**parabolic type**)

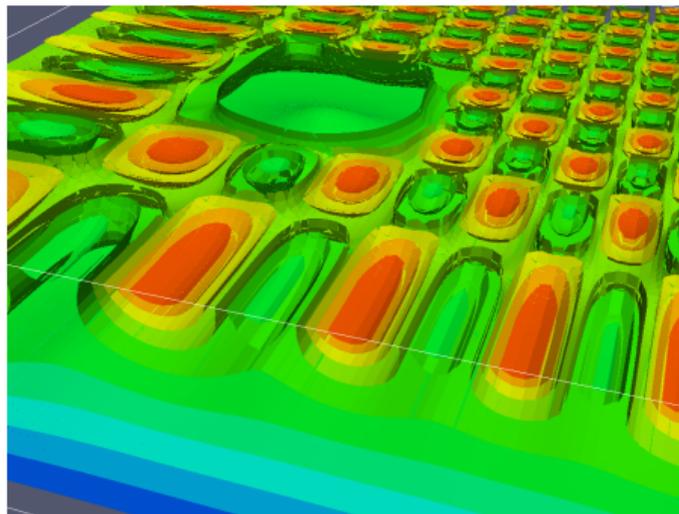
$$\begin{aligned}\partial_t u - \Delta u &= f && \text{in } \Omega \times \Sigma, \Sigma = (t_0, t_0 + T) \\ u &= u_0 && \text{at } t = t_0 \\ u &= g && \text{on } \partial\Omega\end{aligned}$$

- ▶ Wave equation (sound propagation) (**hyperbolic type**)

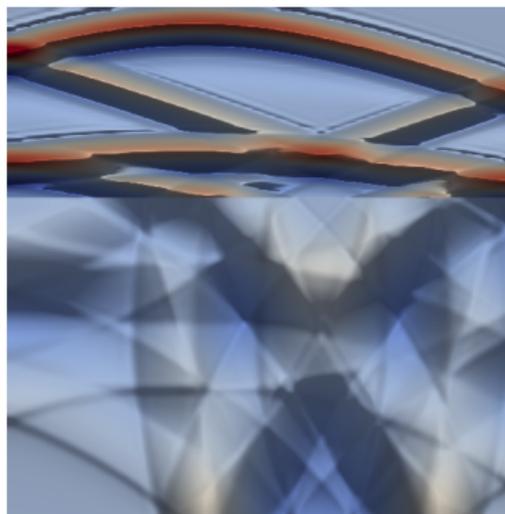
$$\partial_{tt} u - \Delta u = 0 \quad \text{in } \Omega$$

# Second Order Model Problems

Solutions have different behavior



(parabolic)



(hyperbolic)

Preview: The Finite Element Method

# What is a Solution to a PDE?

**Strong form:** Consider the model problem

$$-\Delta u + u = f \quad \text{in } \Omega, \quad \nabla u \cdot \nu = 0 \quad \text{on } \partial\Omega$$

Assume  $u$  is a solution and  $v$  is an arbitrary (smooth) function, then

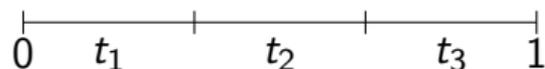
$$\begin{aligned} & \int_{\Omega} (-\Delta u + u)v \, dx = \int_{\Omega} fv \, dx \\ \Leftrightarrow & - \int_{\Omega} (\nabla \cdot \nabla u)v \, dx + \int_{\Omega} uv \, dx = \int_{\Omega} fv \, dx \\ \Leftrightarrow & \int_{\Omega} \nabla u \cdot \nabla v \, dx - \int_{\partial\Omega} (\nabla u \cdot \nu)v \, dx + \int_{\Omega} uv \, dx = \int_{\Omega} fv \, dx \\ \Leftrightarrow & \int_{\Omega} \nabla u \cdot \nabla v + uv \, dx = \int_{\Omega} fv \, dx \\ \Leftrightarrow & a(u, v) = l(v) \end{aligned}$$

**Weak form:** Find  $u \in H^1(\Omega)$  s. t.  $a(u, v) = l(v)$  for all  $v \in H^1(\Omega)$ .

# The Finite Element (FE) Method

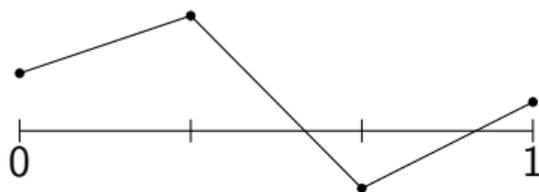
**Idea:** Construct finite-dimensional subspace  $U \subset H^1(\Omega)$

Partition domain  $\Omega$  into “elements”  $t_i$ :



$$\Omega = (0, 1), \quad T_h = \{t_1, t_2, t_3\}$$

Construct function from piecewise polynomials, e.g. linears:



$$U_h = \{u \in C^0(\Omega) : u|_{t_i} \text{ is linear}\}$$

**Insert in weak form:**  $U_h = \text{span}\{\phi_1, \dots, \phi_N\}$ ,  $u_h = \sum_{j=1}^N x_j \phi_j$ , then

$$u_h \in U_h : a(u_h, \phi_i) = l(\phi_i), \quad i = 1, \dots, N \quad \Leftrightarrow \quad \boxed{Ax = b}$$

DUNE Software

- ▶ **Many different PDE applications**
  - ▶ Multi-physics
  - ▶ Multi-scale
  - ▶ Inverse modeling: parameter estimation, optimal control
- ▶ **Many different numerical solution methods, e.g. FE/FV**
  - ▶ No single method to solve all equations!
  - ▶ Different mesh types: mesh generation, mesh refinement
  - ▶ Higher-order approximations (polynomial degree)
  - ▶ Error control and adaptive mesh/degree refinement
  - ▶ Iterative solution of (non-)linear algebraic equations
- ▶ **High-performance Computing**
  - ▶ Single core performance: Often bandwidth limited
  - ▶ Parallelization through domain decomposition
  - ▶ Robustness w.r.t. to mesh size, model parameters, processors
  - ▶ Dynamic load balancing in case of adaptive refinement

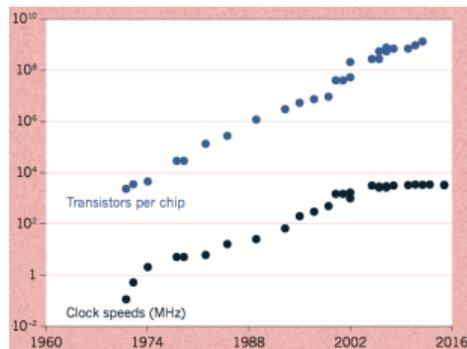
## Distributed and **U**nified **N**umerics **E**nvironment

**Domain specific abstractions for the numerical solution of PDEs with grid based methods.**

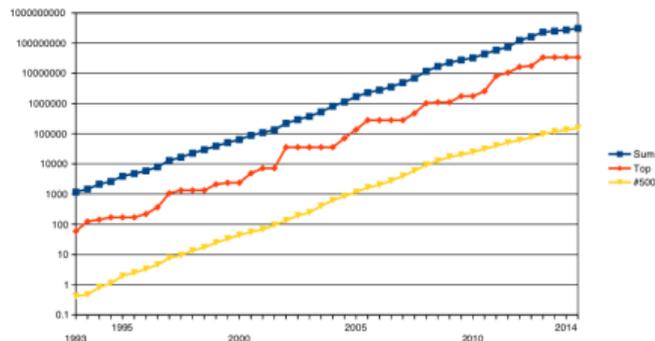
### Goals:

- ▶ Flexibility: Meshes, discretizations, adaptivity, solvers.
- ▶ Efficiency: Pay only for functionality you need.
- ▶ Parallelization.
- ▶ Reuse of existing code.
- ▶ Enable team work through standardized interfaces.

# Trends in Computer Architecture



Moore's law  
Nature, 530 (2016), pp. 144-147



Supercomputer performance (GFLOPs/s)  
<https://commons.wikimedia.org/w/index.php?curid=33540287>

- ▶ Power wall
  - ▶ Power consumption is limiting factor for exascale computing
  - ▶ Clock rate stagnates but Moore's law is still valid
- ▶ Memory wall
  - ▶ Bandwidth not sufficient to sustain peak performance
- ▶ ILP wall
  - ▶ Revival of vectorization in form of SIMD instructions

- ▶ Solution of large sparse algebraic systems  $F(z) = 0$
- ▶ Consider linear case  $Ax = b$ ,  $A \in \mathbb{R}^{N \times N}$ :
  - ▶ Non-sparse Gauß elimination:  $O(N^3)$
  - ▶ Sparse Gauß elimination:  $O\left(N^{\frac{3(d-1)}{d}}\right)$
  - ▶ Multigrid:  $O(N)$

$N$	Gauß elimination $\frac{2}{3}N^3$	multigrid $1000N$
1000	0,66 s	0,001 s
10000	660 s	0,01 s
100000	7,6 days	0,1 s
$10^6$	21 years	1 s
$10^7$	21.000 years	10 s

Run-time @ 1 GFLOPs/s

## AMG Weak Scaling Results

- ▶ AAMG in DUNE is Ph. D. work of **Markus Blatt**
- ▶ BlueGene/P at Jülich Supercomputing Center
- ▶  $P \cdot 80^3$  degrees of freedom ( $5120^3$  finest mesh), CCFV
- ▶ Poisson problem,  $10^{-8}$  reduction
- ▶ AMG used as preconditioner in BiCGStab (2 V-Cycles!)

procs	1/h	lev.	TB	TS	lt	Tlt	TT
1	80	5	19.86	31.91	8	3.989	51.77
8	160	6	27.7	46.4	10	4.64	74.2
64	320	7	74.1	49.3	10	4.93	123
512	640	8	76.91	60.2	12	5.017	137.1
4096	1280	10	81.31	64.45	13	4.958	145.8
32768	2560	11	92.75	65.55	13	5.042	158.3
262144	5120	12	188.5	67.66	13	5.205	256.2

Note: Lecture on Friday, October 22 will be given by Michal Tóth!